

## Some parameters for volcanic conduit eruption models

Requested by Volcanology Workshop, provided by Youxue Zhang

### Bulk H<sub>2</sub>O diffusivity in rhyolitic melt:

$$D = 10^{-12} X \exp(m) \left\{ 1 + \exp \left[ 56 + m + X \left( -34.1 + \frac{44620}{T} + \frac{57.3P}{T} \right) - \sqrt{X} \left( 0.091 + \frac{4.77 \times 10^6}{T^2} \right) \right] \right\}$$

where  $m = -20.79 - 5030/T - 1.4P/T$ ,  $P$  is in MPa,  $T$  is in K,  $X$  is the mole fraction of total H<sub>2</sub>O on a single oxygen basis, and  $D$  is in m<sup>2</sup>/s. The experimental data are for Mono Crater rhyolite. The reference is Zhang and Behrens (2000). To calculate  $X$  from wt% of total H<sub>2</sub>O ( $W$ ), use the following formula for rhyolite:

$$X = \frac{W/18.015}{W/18.015 + (100 - W)/32.49}$$

### H<sub>2</sub>O solubility in rhyolitic melt:

$$W = \sqrt{P} \left( 0.4874 - \frac{608}{T} + \frac{489530}{T^2} \right) + P \left( -0.06062 + \frac{135.6}{T} - \frac{69200}{T^2} \right) + P^{3/2} \left( 0.00253 - \frac{4.154}{T} + \frac{1509}{T^2} \right),$$

where  $P$  is in MPa,  $T$  is in K, and  $W$  is total H<sub>2</sub>O solubility in wt%. That is, when  $W = 1$ , it means that there is 1 wt% total H<sub>2</sub>O. The model is applicable to Mono Crater rhyolite. For reference, "the equation is reparameterized and simplified from the solubility model of Zhang (1999) by Youxue Zhang." A new paper by Liu et al. will be submitted soon and there will be a new solubility model. When the time comes, I will provide an updated equation and reference.

### Bulk H<sub>2</sub>O diffusivity in basaltic melt:

$$D = W \cdot \exp(-10.88 - 15200/T),$$

where  $T$  is in K,  $D$  is in m<sup>2</sup>/s, and  $W$  is wt% of total H<sub>2</sub>O (if there is 1 wt% total H<sub>2</sub>O, then  $W = 1$ ). The reference should be as follows: "The equation is reparameterized from Zhang and Stolper (1991)". The experimental data only cover the range of 0.2 to 0.4 wt% total H<sub>2</sub>O (1300-1500°C and 1.0 GPa). Extrapolation to 1.0 wt% total H<sub>2</sub>O should be OK, to 2.5 wt% total H<sub>2</sub>O might have an uncertainty of a factor of 5. Extrapolations to even higher total H<sub>2</sub>O is not recommended. Clearly more experimental data on H<sub>2</sub>O diffusion in basaltic melt to cover higher H<sub>2</sub>O content and different pressures are necessary.

### CO<sub>2</sub> diffusivity in basaltic melt:

$$D = \exp(-7.96 - 23453/T). \quad (\text{Watson, 1994})$$

The above formula is actually for an anhydrous sodium aluminosilicate melt at 50 MPa but diffusivity in basaltic melt has been shown to be indistinguishable from it (Watson, 1994). CO<sub>2</sub> diffusion in hydrous basaltic melt has not been experimentally studied. Because CO<sub>2</sub> diffusivity seems to be independent of anhydrous melt composition, one might use data for rhyolitic melts. CO<sub>2</sub> diffusivity in rhyolitic melts has been investigated for two H<sub>2</sub>O concentrations, one is dry, and the other with 8.0 wt% H<sub>2</sub>O (Watson, 1994). No general equation is available for CO<sub>2</sub> diffusivity as a function of  $P$ ,  $T$  and H<sub>2</sub>O content.

**Density of basaltic melt as a function of T, P and W:**

Need basaltic melt composition. See Ochs and Lange (1999).

**Surface tension of basaltic melt:**

For dry basaltic melt, surface tension is 0.37 N/m (Walker and Mullin, 1981).

Khitarov et al. (1979) reported the following surface tension values at 1200°C:

P <sub>total</sub> (atm)	H <sub>2</sub> O (wt%)	CO <sub>2</sub> (wt%)	surface tension (N/m)
1			0.4
1000			0.16 to 0.17
3000	4.5	0.3	0.10
5000			0.09

Since the work of Khitarov et al. (1979) was carried out many years ago and at the time measurement of H<sub>2</sub>O and CO<sub>2</sub> was not easy, the quality of the data (such as volatile concentrations) is not known. No attempt is made to make the data self-consistent (e.g., with solubility model).

**References:**

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